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# Performance of Different Models for Curve Number Estimation (Case study: Bar Watershed in Khorasan Razavi Province, Iran)

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ABSTRACT Among different models for runoff estimation in watershed management, the Soil Conservation Services-Curve Number (SCS-CN) method along with its modifications have been widely applied to ungauged watersheds because of quickly and more accurate estimation of surface runoff. This approach has been widely accepted by hydrologists, water resources planners, foresters, and engineers, as well. Therefore, this work was aimed to estimate the curve number using CN-values through several methods viz. SCS, Sobhani (1975), Hawkins *et al.* (1985), Chow *et al.* (1988), Neitsch *et al.* (2002) and Mishra *et al.* (2008) in Bar Watershed, Iran. According to the results, the Neitsch formula showed the best performance for estimating the Curve Number in situation with low (CNI) and high (CNIII) antecedent moisture conditions. However, the weakest performance was related to Mishra (2008) in CNI and CNIII-conversions. The weakest performance was resulted from the exponential form of the Neitsch *et al.* formula and the variable meteorological conditions of the Bar Watershed over the year.

**Key words**: Antecedent soil moisture, Flood estimation, North-Eastern Iran, Rainfall-runoff modeling

#### 1 INTRODUCTION

Nowadays, integrated watershed management (IWM) has an important function in many fields. In fact, it implies correct and appropriate use of water to land and other natural resources in a watershed for runoff estimation which is required for planning, developing and managing water resources. Runoff is one of the significant hydrological variables which is used in the water resources applications and planning

management (Amutha and Porch Elvan, 2009). Relationship between rainfall-runoff is highly nonlinear, complex and complicated and dependent to many factors such as rainfall intensity and duration, soil type, antecedent soil moisture (AMC), land use, evaporation, infiltration, land cover, and slope (Elhakeem and Papanicolaou, 2009). Hydrologic models have been widely used to explain and predict complex behaviors associated with the

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management of environmental systems (Schulze, 2000; Bronstert et al., 2002; Croke et al., 2004; Siriwardena et al., 2006; Lin et al., 2007; Kalin and Hantush, 2009; Isik, et al., 2012). In this way, there are many hydrologic models which are used to estimate runoff. However, physically based models are faced with some limitations because of their large number of input parameters and complicated calibration (Wu et al., 1993; Kothyari and Jain, 1997; Xiao et al., 2011). Regarding the numerous variables and uncertainties governing the rainfall-runoff process, the lumped-conceptual models are useful approaches for hydrological analysis (Ponce and Hawkins, 1996; McCuen, 2003; Mishra and Singh, 2003; Elhakeem and Papanicolaou, 2009). However, these models must be calibrated by using measurements (Papanicolaou et al., 2008). Due to the simplicity, high speed computing, correct estimation, and few data requirement, the Soil Conservation Service method (SCS) curve number (CN) is one of the best methods in small agricultural and urban watersheds (Ponce and Hawkins, 1996; Bhuyan et al., 2003; Liu and Li, 2008; Xiao et al., 2011). Additionally, many papers revealed that curve number has been incorporated into a wide range of single event and continuous computer models (Ponce, 1989). The SCS curve number is a function of the soils ability to allow infiltration of water with respect to the land use, land cover and AMC. According to the U.S. Soil Conservation Service classification, there are four hydrologic soil groups, i.e. A, B, C and D with respect to the rate of runoff potential and final infiltration rate (Amutha and Porch Elvan, 2009). Actually, this method is a dual parameter model for predicting surface runoff depth from rainfall depth of individual storm events and it has been widely accepted by scientists, resources hydrologists, water planners,

agronomists, foresters, and engineers for surface runoff estimation (Patil *et al.*, 2008).

The SCS-CN (method SCS 1956, 1964, 1971, 1993) transforms rainfall to surface runoff (or rainfall-excess) by using curve number, which is derived from watershed characteristics and antecedent 5-day rainfall (Mishra et al., 2008). Since a natural watershed is very dynamic and has different reaction versus storms, some parameters such as antecedent 5-day rainfall, interception and soil moisture show a variety in individual CN of a watershed even during a storm. Hjelmfelt (1980) has shown that the curve number equation is identical for the special case with the constant rainfall intensity and zero asymptotic infiltration rates. Moreover, this method has been used in ungauged rural watersheds and has been evolved well beyond its original objective for surface runoff prediction in urbanized and forested watersheds (USDA, 1986).

Some scientists such as Ponce (1989) believe that the CN method should not apply for the watersheds which are longer than 250 km<sup>2</sup> and are not subdivided. However, sometimes this method is used for the situation that is not applicable appropriate (Suresh Babu et al., 2008; Xiao et al., 2011). Although this method has many advantages, some parameters such as spatial and temporal infiltration and time distribution are not considered in that. Furthermore, CNs are computed, empirically, but other factors such as soil and vegetation which effect on them are not empirically computed. The division of soils into hydraulic groups is very coarse and the definition of the antecedent moisture condition is not a quantitative variable (Ponce and Hawkins, 1996; Xiao et al., 2011). Bhuyan et al. (2003) has studied event based watershed scale of the AMC values to adjust field-scale CNs as well as to identify the hydrologic parameters that would provide the best estimate of the AMC. This study showed that the AGNPS (Agricultural Non-Point Source Model) watershed model (Heaney et al., 2001) overestimated the runoff depth while using a CN based on AMC-II condition (Patil et al., 2008; El-Hames, 2012). Thus, the universal assumption for applying AMC-II conditions under typical watershed conditions was observed to be invalid for many experimental watersheds (Patil et al., 2008). Xiao et al. (2011) applied the SCS-CN model to estimate runoff in Loess Plateau of China, as well. Carlesso et al. (2011) wanted to measure the runoff for different soils classes at different rainfall intensities in Southern Brazil. For this aim each class of soil, the initial time and runoff rate, rainfall characteristics, surface slope, crop residue amount and cover percentage, soil densities, soil porosity, textural fractions, and the initial and saturated soil water content were measured. The runoff measured was compared to Smith's modified and Curve Number (USDA-SCS) models. The Smith's modified model overestimated the cumulative runoff by about 4%. The SCS Curve Number model overestimated the cumulative runoff by about 34%. Smith's modified model better estimated the surface runoff for soil with high soil water content, and it was considered satisfactory for Southern Brazil runoff estimations. The SCS Curve Number model overestimated the cumulative runoff and its use needs adjustments particularly for no-tillage management system.

Deshmukh et al. (2013) used three different methods for three watersheds located in Narmada basin. In this work, the CN computed from the observed rainfallrunoff events was termed as CN (PQ), land use and land cover (LULC) was termed as CN (LU) and the CN based on land slope was termed as SACN2. The estimated annual CN (PQ) varied from 69 to 87 over the 26 years period with a median of 74 and an average of 75. The CN (PQ) ranged from 70 to 79 were the most significant values and truly represented the AMC-II condition for the Sher Watershed. The annual CN (LU) was computed for all three watersheds using GIS for the years 1973, 1989 and 2000. The computed CN (LU) values increasing trend with the time which was attributed to the expansion of agricultural area in whole watersheds. The predicted values of CN (LU) used to predict runoff potential under the LULC alterations. Comparison of CN (LU) with CN (PQ) values showed a close agreement which establish the validation of the LULC classification. In addition, results showed that for the micro watershed planning, SCSmethod should be modified to incorporate the effect of change in land use and land cover along with land slope alteration. Lin et al. (2014) combined Xinanjiang model with the Curve Number (XAJ-CN model) to simulate the impact of land use change on water flow in Dongjiang River basin which constitutes the most important water source system Guangdong Province and Hong Kong City. They calibrated and validated the model based on the 10 years data. It was observed that the simulated runoff matched the observed one, which indicated that the performance of the XAJ-CN model was satisfactory. Results showed that the impact of land use change on runoff was more obvious flood season compared to that in dry season. The impact of changes in the CN value on surface runoff was the highest flood season, while the change in the CN value

mainly affected groundwater recharge dry season. Lal et al. (2015) used a large number of observed rainfall and runoff events occurred in India. The effects of land features such as slope and antecedent moisture was evaluated on the curve number parameters of the SCS-CN methodology. Results showed that the original assumption of the optimized initial abstraction ratio ( $\lambda$ ) of 0.20 was unusually high. The median and mean  $\lambda$ values were respectively 0 and 0.034 for natural rainfall and runoff data and 0.033 and 0.108 for ordered rainfall and runoff data, respectively. In addition, CN or potential maximum retention (S) values showed a degree of dependence on the physically observed 1-day antecedent soil moisture than other duration antecedent soil moisture values. The AMC was introduced as the initial moisture condition of the watershed before storm event. Commonly, the AMC-II is considered as the base condition for CN determining.

Curve numbers was first calculated under the AMC-II condition, then they were adjusted based on the AMC-III or AMC-I depending on the 5-day antecedent amount of rainfall (Mishra et al., 2008). There were three drawbacks with this assumption (Hope and Schulze, 1981). First, the relationship between antecedent rainfall and AMC was defined for discrete classes, rather than continuous (Hawkins, 1978; Mishra et al., 2008). Second, the use of 5-day antecedent rainfall was more applicable on subjective judgments than physical reality. Third, evapotranspiration and drainage were not considered in depletion of moisture (Mishra et al., 2008). After determination of the AMC and CN-values by NEH-41 (SCS, 1972), some scientists such as Sobhani (1975), Hawkins et al. (1985) and Chow et al. (1988) have shown about the other formulae same CNconversion. Neitsch et al. (2002) and Mishra et al. (2008) have also represented CNconversion formulae entirely different in form and these are being used in the Soil and Water Assessment Tools (SWAT). Therefore, there are many formulae about this subject which are very different in terms of the techniques for CN estimation. While runoff estimation is very important factor for watershed management and planning; it is quite necessary to compare the mentioned conversion formulae for CN estimation and discuss their validity with regard to some evaluation criteria, which is the main purpose of this work.

#### 2 MATERIALS AND METHODS

#### 2.1 Study Area

The Bar Watershed is located in Khorasan Ravazi Province, northeastern Iran. The area of this watershed is about 111 km² which is located between 36° 27′ 38″ to 36° 36′ 32″ N-latitude and 58° 40′ 46″ to 58° 49′ 31″ Elongitude. The mean annual rainfall and average altitude is about 330 mm and 2226 m, respectively with a semi-arid climate (Jafari *et al.*, 2012). The length of Bar main river is 22.5 km and its average slope is 4.2% which finally drains to the Neyshabur plain (Sadeghi *et al.*, 2010). Figure 1 shows a general view of the Bar Watershed.

<sup>1.</sup> National Engineering Handbook

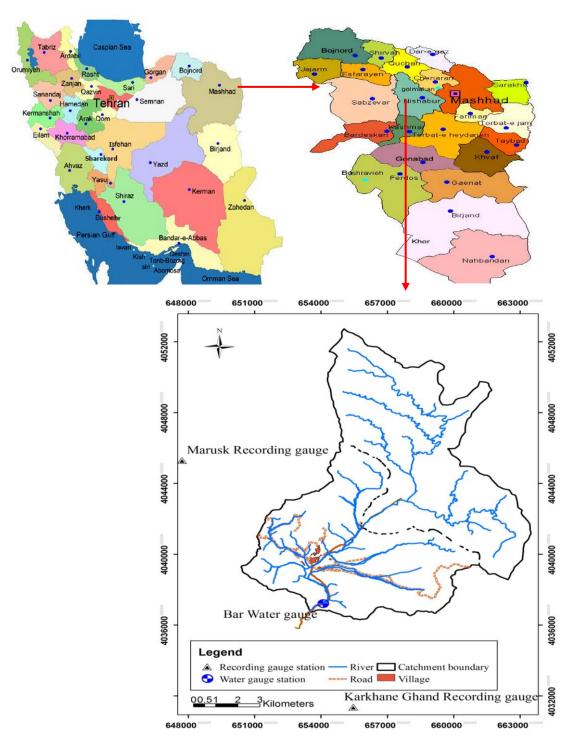


Figure 1 The study area location, rainfall stations and gauge site

#### 3 METHODOLOGY

In order to conduct the study, daily discharge and precipitation data during the period 1951-2006

(56 years) was used. According to Xiao *et al.* (2011) and many other researcher's observations; by increasing the initial abstraction, the ratio of

rainfall ( $I_a/S$ ) increased gradually when it is lower than 50 mm, but when it is more than 50 mm, the  $I_a/S$  value increased rapidly. The popular form of the SCS-CN equation obtained through combining water balance and two fundamental hypotheses has been represented as the SCS approach (1972):

$$Q = \frac{\left(P - Ia\right)^2}{P - Ia + S} \tag{1}$$

Where P is the total rainfall in mm, Q is the direct runoff in mm, S is the potential maximum retention and Ia is the initial abstraction which involve the interception, surface storage and initial infiltration and many other factors so that it is expressed as a function of S (Ia= $\lambda$ S). In this equation,  $\lambda$  is initial abstraction coefficient which depends on geological and climatic factor and it is between 0.1 and 0.3. However, many books and papers empirically assume  $\lambda$  is 0.2 (SCS, 1985). From the observed rainfall—runoff data and the SCS-CN parameter, S can be determined by solving Eq. 1 for  $\lambda$  = 0.2, as follows (Hawkins, 1978):

$$S = 5 \left[ (P + 2Q) - \sqrt{Q(4Q + 5P)} \right]$$
 (2)

Where S is related to the curve number (CN) which it is very variable from 0 up to 100. If it is 0, there is not any direct runoff and therefore 100 represents that all rainfalls turn into runoff. CN = 100 represents a theoretical lower bound of the potential retention storage, and CN = 0 denotes a theoretical upper bound of the potential retention storage (Jung *et al.*, 2012). The relation between two parameters is shown as:

$$CN = \frac{25400}{S + 254} \tag{3}$$

Where S is in mm and CN is a non-dimensional factor and shows runoff potential which is controlled by the AMC, land use, soil type, and treatment (SCS, 1985). There are three antecedent moisture conditions (AMC), i.e. AMCI, AMCII and AMCIII which are defined in dry, medium or normal and wet soil conditions. In this formula median CN (CNII) selected as a representative CN which is valid for normal antecedent moisture condition of the watershed (Xiao *et al.*, 2011). This 'Median CN' approach is commonly adopted (Hjelmfelt, 1991; Hawkins *et al.*, 2002; Mishra *et al.*, 2005; Jung *et al.*, 2012).

In this work, runoff depth was calculated for every day in all of the months during 56 years through converting discharge to volume flood in Excel 2010. Then, for every daily runoff, potential maximum retention (S) and curve number (CN) was computed daily by using equations 2 and 3. It is noticeable that the curve number was defined in accordance with CNII which was the watershed's "average condition" in terms of wetness by using Table 1 and then adjusting to AMC III or AMC I depending on the 5-day depth of antecedent rainfall. This table belongs to the NEH-4 CNvalues (SCS, 1972). In order to compare the performance of the discussed CN methods with the observed data from a gauged watershed; the recorded data at the watershed outlet were analyzed to obtain the CN with different methods by using Tables 1 and 2. These observed runoff depths of different rainfall events were utilized to compare with the CN based predictions to analyze performances. All proposed conversion formulae were represented in Table 2. Sobhani (1975), Hawkins et al. (1985), Chow et al. (1988), Nietsch et al. (2002), and Mishra et al. (2008) formulas were used for comparing their ability through three parameters indices for evaluation criteria viz. RE, RMSE and R<sup>2</sup>.

**Table 1** Curve number (CN) and constants for the case Ia=0.2 S (SCS, 1972)

| 1               | 2              | 3                | 4              | 5                              | 1               | 2              | 3                | 4                | 5                                   |
|-----------------|----------------|------------------|----------------|--------------------------------|-----------------|----------------|------------------|------------------|-------------------------------------|
| CN for<br>AMCII | CN for<br>AMCI | CN for<br>AMCIII | S<br>(mm)      | Curve<br>starts<br>where<br>P= | CN for<br>AMCII | CN for<br>AMCI | CN for<br>AMCIII | S<br>(mm)        | Curve<br>starts<br>where P=<br>(mm) |
| 100             | 100            | 100              | 0              | ( <b>mm</b> )                  | 60              | 40             | 78               | 169.41           | 33.78                               |
| 99              | 97             | 100              | 2.56           | 5.08                           | 59              | 39             | 78<br>77         | 167.38           | 35.30                               |
| 98              | 94             | 99               | 5.18           | 10.16                          | 58              | 38             | 7 <i>7</i><br>76 | 183.89           | 36.83                               |
| 98<br>97        | 9 <del>1</del> | 99               | 7.84           | 15.24                          | 57              | 37             | 75               | 191.51           | 38.35                               |
| 96              | 89             | 99               | 10.59          | 20.32                          | 56              | 36             | 75               | 191.51           | 39.87                               |
| 95              | 87             | 98               | 13.36          | 20.32                          | 55              | 35             | 73<br>74         | 207.77           | 41.65                               |
| 93<br>94        | 85             | 98               | 16.20          | 3.30                           | 54              | 34             | 73               | 216.40           | 43.18                               |
| 93              | 83             | 98               | 19.12          | 3.81                           | 53              | 33             | 73               | 225.29           | 44.95                               |
| 93<br>92        | 81             | 97               | 22.09          | 4.31                           | 52              | 32             | 72               | 234.44           | 46.99                               |
| 91              | 80             | 97<br>97         | 25.12          | 5.08                           | 51              | 31             | 70               | 244.09           | 48.76                               |
| 90              | 78             | 96               | 28.19          | 5.58                           | 50              | 31             | 70               | 254.00           | 50.80                               |
| 89              | 76<br>76       | 96               | 31.49          | 6.35                           | 49              | 30             | 69               | 264.16           | 52.83                               |
| 88              | 75             | 95               | 34.54          | 6.85                           | 48              | 29             | 68               | 274.32           | 54.86                               |
| 87              | 73             | 95<br>95         | 37.84          | 7.62                           | 47              | 28             | 67               | 287.02           | 57.40                               |
| 86              | 73             | 93<br>94         | 41.40          | 8.38                           | 46              | 27             | 66               | 297.18           | 59.43                               |
| 85              | 70             | 94<br>94         | 44.70          | 8.89                           | 45              | 26             | 65               | 309.88           | 59.43<br>61.97                      |
| 84              | 68             | 93               | 48.26          | 9.65                           | 43              | 25             | 64               | 322.58           | 64.51                               |
| 83              | 67             | 93               | 52.07          | 10.41                          | 43              | 25<br>25       | 63               | 335.28           | 67.05                               |
| 82              | 66             | 93<br>92         | 55.88          | 10.41                          | 43              | 23             | 62               | 350.52           | 70.10                               |
| 81              | 64             | 92<br>92         | 59.43          | 11.17                          | 42              | 23             | 61               | 365.76           | 73.15                               |
| 80              | 63             | 92<br>91         | 63.5           | 12.70                          | 40              | 22             | 60               | 381.00           | 76.20                               |
| 79              | 62             | 91               | 67.56          | 13.46                          | 39              | 21             | 59               | 396.24           | 79.24                               |
| 79<br>78        | 60             | 90               | 71.62          | 14.22                          | 38              | 21             | 58               | 414.02           | 82.80                               |
| 78<br>77        | 59             | 90<br>89         | 71.02<br>75.94 | 15.24                          | 38<br>37        | 20             | 56<br>57         | 431.8            | 86.36                               |
| 76              | 58             | 89               | 80.26          | 16.00                          | 36              | 19             | 56               | 452.12           | 90.42                               |
| 75              | 57             | 88               | 84.58          | 17.018                         | 35              | 18             | 55               | 472.44           | 90.42                               |
| 73<br>74        | 57<br>55       | 88               | 89.15          | 17.018                         | 33              | 18             | 54               | 492.76           | 98.55                               |
| 73              | 53<br>54       | 87               | 93.98          | 18.79                          | 33              | 17             | 53               | 515.62           | 103.12                              |
| 73<br>72        | 53             | 86               | 93.98          | 19.81                          | 33              | 16             | 52               | 513.02           | 103.12                              |
| 72              | 52             |                  | 103.63         | 20.82                          | 31              | 16             | 51               |                  |                                     |
| 70              | 51             | 86<br>85         | 103.03         | 20.82                          | 30              | 15             | 50               | 563.88<br>591.82 | 112.77<br>118.36                    |
| 69              | 50             | 83<br>84         |                |                                |                 | 13             | 43               | 762.00           |                                     |
|                 | 48             |                  | 114.04         | 22.86                          | 25<br>2         |                |                  |                  | 152.40                              |
| 68              |                | 84               | 119.38         | 23.87                          |                 | 9              | 37               | 1016.00          | 203.20                              |
| 67<br>66        | 47<br>46       | 83               | 124.96         | 24.89                          | 15              | 6              | 30               | 1440.18          | 288.04                              |
| 66<br>65        | 46<br>45       | 82<br>82         | 130.81         | 26.16                          | 10              | 4              | 22               | 2286.00          | 457.20                              |
| 65<br>64        | 45             | 82               | 136.65         | 27.43                          | 5               | 2              | 13               | 4826.00          | 965.20                              |
| 64              | 44             | 81               | 142.74         | 28.44                          | 0               | 0              | 0                | infinity         | infinity                            |
| 63              | 43             | 80               | 149.09         | 29.71                          |                 |                |                  |                  |                                     |
| 62              | 42             | 79               | 155.70         | 31.24                          |                 |                |                  |                  |                                     |
| 61              | 41             | 78               | 162.30         | 32.51                          |                 |                |                  |                  |                                     |

Table 2 Some methods about CN-conversion formulae

| Methods                      | AMC I   | AMC III   |
|------------------------------|---|---|
| Sobhani (1975)               | $CN_{II} = \frac{CN_{II}}{2.334 - 0.0133CN_{II}}$   | $CN_{III} = \frac{CN_{II}}{0.4036 + 0.0059CN_{II}}$   |
| Hawkins <i>et al.</i> (1985) | $CN_{I} = \frac{CN_{II}}{2.281 - 0.0128CN_{II}}$  | $CN_{III} = \frac{CN_{II}}{0.427 + 0.0057CN_{II}}$    |
| Chow <i>et al.</i> (1988)    | $CN_{I} = \frac{4.2CN_{II}}{10 - 0.058 \ CN_{II}}$  | $CN_{III} = \frac{23CN_{II}}{10 + 0.13CN_{II}}$       |
| Neitsch <i>et al.</i> (2002) | $CN_{I} = CN_{II} - \frac{20 (100 - CN_{II})}{\{100 - CN_{II} + exp \left[ 2.533 - 0.063 (100 - CN_{II}) \right] \}}$ | $CN_{III} = CN_{II} exp \{ 0.0067 (100 - CN_{II}) \}$ |
| Mishra <i>et al.</i> (2008)  | $CN_{II} = \frac{CN_{II}}{2.2754 - 0.0128CN_{II}}$  | $CN_{III} = \frac{CN_{II}}{0.430 + 0.057CN_{II}}$     |

Root mean square error (RMSE), relative error (RE), and Pearson correlation coefficient ( $R^2$ ) are defined as follows:

$$RMSE = \sqrt{\frac{\sum_{i=1}^{N} (Q_{O} - Q_{e})^{2}}{N}}$$
 (5)

$$RE = \left| \frac{Qo - Qe}{Qo} \right| *100 \tag{6}$$

$$R^{2} = 1 - \frac{\sum_{i=1}^{n} (Qo - Qe)^{2}}{\sum_{i=1}^{n} (Qo - \overline{QO})^{2}}$$
 (7)

Where RMSE and RE are two evaluation criteria for indicating difference between the model simulation and observation,  $Q_o$  is observed flow,  $Q_e$  is simulated flow and N is the number of data records. RE is expressed in percent (%). In fact CN in NEH-4 (Table 2) is an observation and those derived from the above formulae are simulated, computed and presented in Table 2. So that, for daily runoff CNI and CNIII was computed with every method which was written in Table 1. Results are shown in Tables 3 and 4, respectively.

### 4 RESULTS AND DISCUSSIONS

As it was explained, CN in NEH-4 is an average of CNs and it is defined as the observation data, as well. For the period 1951 to 2006, the value of CNII was calculated and transformed to CNI and CNIII by using NEH-4 formula (SCS, 1972) as the target values

(56 common years) and after that Sobhani (1975), Hawkins *et al.* (1985), Chow *et al.* (1988), Nietsch *et al.* (2002), and Mishra *et al.* (2008) formulas were applied (Table 1). It is noticeable that, undesirable negative yield values of CNI in CNII results in negative S-values which are not conceptually rational indeed. For this reason, these results have been omitted from other results. In addition, when precipitation was less than discharge, the results was omitted, as well.

Since there is a record of data (about 20300) the average of relative CNI, II and III were written and therefore average of relative error, root mean square and correlation coefficient were presented in Tables 3 and 4. As it can be seen in Table 3, the lowest and the highest values of CNI in NEH-4 (observed data) was 77.94 and 92.29 in April and July, respectively while in computed data the lowest and the highest value of CNI was 78.41 and 93.12 in these months, in Neitsch et al. (2002) and Mishra et al. (2008) formulas respectively. Therefore, the range of observed data CNI in NEH-4 varied from 77.94 to 92.29 and in computed data from 78.41 to 93.12. Additionally, the lowest and the highest values of observed data CNIII in NEH-4 was 95.32 and 98.60 in April and July

respectively while in computed data the lowest and the highest values of CNI were 90.15 and 98.60 in September and July in Mishra *et al.* (2008) and Neitsch *et al.* (2002) formulas respectively. Thus, the range of observed data CNIII in NEH-4 varied from 95.32 to 98.60 and from 90.15 to 98.60 in computed data.

Furthermore, the range of average relative error for CNI was (0.76-1.75), (0.92-2), (0.69-1.53), (0.48-0.82), and (0.93-2.02) and for CNIII was (0.21-0.47), (0.28-0.62), (0.31-0.71), (0.20-0.28) and (0.29-7.74) that were belonged to Sobhani (1975), Hawkins et al. (1985), Chow et al. (1988), Nietsch et al. (2002) and Mishra et al. (2008) formulas, respectively. According to Table 3, the lowest and the highest relative errors for CNI and CNIII belonged to Neitsch et al. (2002) and Mishra et al. (2008), in respective. In fact the Neitsch et al. (2002) formulae exhibited the narrowest range of RE-variation and as result it was the closest one to NEH-4 data (observation data). Therefore among all formulas, Mishra et al. (2008) and Neitsch et al. (2002) were the worst and the best models in this study which their results did not agree with Mishra et al. (2008) as can be seen clearly in Table 4.

Table 3 Comparison of average of various AMC dependent CN and their respective RE

|      | Z     | NEH-4    |       | S        | Sobhani (1975) | 1975) |        | Hav   | Hawkins <i>et al.</i> (1985) | al. (19 | <b>)</b> 85) | C     | Chow et al. (1988) | l. (198 | 8)    | Ni        | Nietsch <i>et</i> | et al. (2002) | 02)    | Mi    | Mishra <i>et al.</i> (2008) | al. (20 | 108)   |
|------|-------|----------|-------|----------|----------------|-------|--------|-------|------------------------------|---------|--------------|-------|--------------------|---------|-------|-----------|-------------------|---------------|--------|-------|-----------------------------|---------|--------|
|      | Obs   | Observed |       | Computed | uted           | RE    | RE (%) | Com   | Computed                     | RE (%)  | (%)          | Com   | Computed           | RE (%)  | (%)   | Com       | Computed          | RE            | RE (%) | Com   | Computed                    | RE      | RE (%) |
|      | CNII  | CNI      | CNIII | CNI      | CNIII          | CNI   | CNIII  | CNI   | CNIII                        | CNI     | CNIII        | CNI   | CNIII              | CNI     | CNIII | CNI       | CNIII             | CNI           | CNIII  | CNI   | CNIII                       | CNI     | CNIII  |
| Jan  | 91.53 | 82.37    | 96.49 | 83.56    | 96.19          | 1.66  | 0.46   | 83.83 | 95.99                        | 1.91    | 0.61         | 83.32 | 95.92              | 1.50    | 0.67  | 82.83     | 96.55             | 0.78          | 0.28   | 83.86 | 83.86 95.96                 | 1.94    | 0.64   |
| Feb  | 91.06 | 81.80    | 96.22 | 82.93    | 95.94          | 1.53  | 0.38   | 83.20 | 95.73                        | 1.76    | 0.54         | 82.69 | 95.66              | 1.43    | 0.61  | 82.30     | 96.29             | 0.76          | 0.24   | 83.20 | 83.20 95.70                 | 1.74    | 0.56   |
| Mar  | 91.43 | 82.53    | 96.43 | 83.56    | 96.12          | 1.49  | 0.39   | 83.83 | 83.83 95.92                  | 1.73    | 0.55         | 83.33 | 95.85              | 1.34    | 0.61  | 82.94 96. | 47                | 0.72          | 0.25   | 83.83 | 83.83 95.89                 | 1.70    | 0.57   |
| Apr  | 89.01 | 77.94    | 95.32 | 79.10    | 95.01          | 1.63  | 0.41   | 79.43 | 94.75                        | 1.90    | 0.62         | 78.81 | 78.81 94.66        | 1.51    | 0.71  | 78.41     | 95.44             | 0.76          | 0.28   | 79.47 | 79.47 94.71                 | 1.94    | 0.66   |
| May  | 93.31 | 85.57    | 97.30 | 86.81    | 97.02          | 1.53  | 0.34   | 87.03 | 96.86                        | 1.74    | 0.47         | 86.61 | 86.61 96.81        | 1.38    | 0.52  | 86.15     | 97.33             | 0.74          | 0.21   | 87.06 | 87.06 96.84 1.77            | 1.77    | 0.49   |
| Jun  | 95.04 | 89.05    | 98.06 | 90.19    | 97.79          | 1.37  | 0.32   | 90.36 | 90.36 97.67                  | 1.52    | 0.41         | 90.04 | 97.63              | 1.26    | 0.45  | 89.64     | 98.02             | 0.71          | 0.21   | 90.38 | 90.38 97.66                 | 1.55    | 0.43   |
| Jul  | 96.46 | 92.29    | 98.60 | 92.99    | 98.44          | 0.76  | 0.21   | 93.11 | 98.35                        | 0.92    | 0.28         | 92.88 | 98.32              | 0.69    | 0.31  | 92.69     | 98.60             | 0.48          | 0.20   | 93.12 | 98.34                       | 0.93    | 0.29   |
| Aug. | 95.04 | 87.95    | 98.36 | 89.45    | 97.90          | 1.75  | 0.47   | 89.65 | 89.65 97.78                  | 2.00    | 0.59         | 89.27 | 89.27 97.75        | 1.53    | 0.62  | 88.53     | 98.19             | 0.74          | 0.24   | 89.67 | 89.67 97.77                 | 2.02    | 0.60   |
| Sep  | 93.46 | 85.46    | 97.56 | 86.66    | 97.18          | 1.60  | 0.42   | 86.90 | 86.90 97.02                  | 1.90    | 0.57         | 86.45 | 96.97              | 1.34    | 0.63  | 85.84 97  | 97.53             | 0.82          | 0.21   | 86.92 | 86.92 90.15                 | 1.93    | 7.74   |
| Oct  | 93.30 | 85.29    | 97.37 | 86.59    | 97.06          | 1.59  | 0.35   | 86.82 | 86.82 96.90                  | 1.81    | 0.50         | 86.39 | 86.39 96.84        | 1.41    | 0.55  | 85.88     | 97                | .39 0.77      | 0.28   | 86.85 | 96.88                       | 1.86    | 0.52   |
| Nov  | 93.37 | 85.85    | 97.29 | 87.00    | 97.03          | 1.49  | 0.38   | 87.22 | 96.87                        | 1.67    | 0.47         | 86.81 | 96.82              | 1.34    | 0.52  | 86.32     | 97.32             | 0.62          | 0.25   | 87.25 | 96.85                       | 1.69    | 0.49   |
| Dec  | 92.21 | 83.54    | 96.86 | 84.68    | 96.54          | 1.53  | 0.37   | 84.94 | 84.94 96.36                  | 1.80    | 0.54         | 84.46 | 96.29              | 1.37    | 0.61  | 84.00 96. | 96.90             | 90 0.73       | 0.24   | 84.97 | 84.97 96.33                 | 1.83    | 0.57   |

Table 4 Average of RMSE and R2 values for different AMC dependent CN conversion formulae taking SCS (1972, 1985) CN value as target values

|      |      | Sobhar | Sobhani (1975) |           | Ha   | Hawkins <i>et al.</i> (1985) | t al. (19) | (58       | C    | Chow et al. (1988) | al. (198 | 3)        | Ŋ    | Nietsch <i>et al.</i> (2002) | al. (200       | (2)   | M    | Mishra <i>et al.</i> (2008) | al. (200 | 8)        |
|------|------|--------|----------------|-----------|------|------------------------------|------------|-----------|------|--------------------|----------|-----------|------|------------------------------|----------------|-------|------|-----------------------------|----------|-----------|
|      | R    | RMSE   | R              | ${f R}^2$ | RIV  | RMSE                         | R          | ${f R}^2$ | RIV. | RMSE               | R        | ${f R}^2$ | RMSE | SE                           | $\mathbb{R}^2$ | 7,    | RM   | RMSE                        | R        | ${f R}^2$ |
|      | CNI  | CNIII  | CNI            | CNIII     | CNI  | CNIII                        | CNI        | CNIII     | CNI  | CNIII              | CNI      | CNIII     | CNI  | CNIII                        | CNI            | CNIII | CNI  | CNIII                       | CNI      | CNIII     |
| Jan  | 1.62 | 0.73   | 0.998          | 0.995     | 1.81 | 0.84                         | 0.998      | 0.995     | 1.49 | 0.89               | 0.998    | 0.955     | 0.99 | 0.64                         | 0.998          | 0.998 | 1.83 | 0.86                        | 0.998    | 0.995     |
| Feb  | 1.51 | 0.43   | 0.996          | 0.994     | 1.70 | 09.0                         | 0.996      | 0.994     | 1.40 | 0.67               | 966.0    | 0.994     | 06:0 | 0:30                         | 0.999          | 0.997 | 1.70 | 0.63                        | 0.996    | 0.997     |
| Mar  | 1.46 | 0.46   | 0.995          | 0.994     | 1.64 | 0.63                         | 0.995      | 0.994     | 1.35 | 0.70               | 966.0    | 0.994     | 0.87 | 0.32                         | 0.998          | 0.997 | 1.64 | 99.0                        | 0.995    | 0.994     |
| Apr  | 1.51 | 0.46   | 0.998          | 0.997     | 1.73 | 99.0                         | 0.998      | 0.997     | 1.37 | 0.75               | 0.998    | 0.997     | 0.79 | 0.33                         | 0.999          | 966.0 | 1.75 | 69.0                        | 0.998    | 0.997     |
| May  | 1.49 | 0.39   | 0.992          | 0.967     | 1.67 | 0.55                         | 0.992      | 0.967     | 1.35 | 0.61               | 0.993    | 0.967     | 0.78 | 0.26                         | 0.997          | 0.978 | 1.70 | 0.57                        | 0.992    | 0.967     |
| Jun  | 1.41 | 0.38   | 0.992          | 0.991     | 1.57 | 0.49                         | 0.992      | 0.991     | 1.28 | 0.54               | 0.992    | 0.991     | 0.77 | 0.24                         | 0.994          | 0.992 | 1.59 | 0.51                        | 0.992    | 0.985     |
| Jul  | 0.79 | 0.22   | 966.0          | 0.992     | 0.89 | 0.34                         | 0.996      | 0.992     | 0.74 | 0.39               | 0.997    | 0.991     | 0.50 | 0.21                         | 0.999          | 0.998 | 0.90 | 0.36                        | 0.996    | 0.991     |
| Aug. | 1.66 | 0.55   | 966.0          | 0.992     | 1.88 | 0.67                         | 0.995      | 0.992     | 1.48 | 0.72               | 966.0    | 0.992     | 0.73 | 0.28                         | 0.999          | 966.0 | 1.90 | 69.0                        | 0.995    | 0.992     |
| Sep  | 1.53 | 0.46   | 0.995          | 0.993     | 1.74 | 0.62                         | 0.994      | 0.993     | 1.37 | 0.68               | 0.995    | 0.993     | 0.97 | 0.26                         | 0.997          | 966.0 | 1.76 | 9.34                        | 0.995    | 0.993     |
| Oct  | 1.51 | 0.42   | 0.994          | 0.975     | 1.71 | 0.59                         | 0.994      | 0.975     | 1.37 | 0.65               | 0.994    | 0.975     | 0.81 | 0.27                         | 0.996          | 0.979 | 1.73 | 0.61                        | 0.994    | 0.975     |
| Nov  | 1.41 | 0.43   | 0.996          | 0.996     | 1.62 | 0.54                         | 0.996      | 0.995     | 1.25 | 0.59               | 0.996    | 0.995     | 0.61 | 0.31                         | 0.998          | 0.998 | 1.65 | 0.56                        | 0.996    | 0.995     |
| Dec  | 1.49 | 0.43   | 0.996          | 0.994     | 1.70 | 0.61                         | 0.995      | 0.994     | 1.35 | 0.68               | 0.996    | 0.996     | 0.85 | 0.29                         | 0.998          | 0.997 | 1.72 | 0.63                        | 0.995    | 0.994     |

One of the evaluation criteria was RMSE. It is noticeable that the NEH-4 AMC and the Neitsch et al. (2002) definition Tables are not the same which was used to determine soil moisture content of a day as the latter adjusts CNs for AMCs. In the NEH-4 procedure, the season was considered as growing season. In Table 4 all the five methods compared based on average RMSE values which were derived from their application to P-Q data sets of the Bar Watershed during 56 years. These methods can be ranked as follows: The range of average root mean square for CNI was (0.79-1.66), (0.89-1.88), (0.74-1.49), (0.5-0.99) and (0.90-0.19) and for CNIII was (0.22-0.73), (0.34-0.84), (0.39-0.89), (0.21-0.64) and (0.36-9.34) that belonged to Sobhani (1975), Hawkins et al. (1985), Chow et al. (1988), Nietsch et al. (2002), and Mishra et al. (2008) formulas, respectively.

Generally these rankings were the performance of various methods as following (average of RE and RMSE-based):

(For CNII to CNI)

Neitsch> Chow > Sobhani > Hawkins > Mishra

(For CNII to CNIII)

Neitsch > Sobhani > Hawkins > Chow > Mishra

Since the number of figures about correlation coefficient are too many (N=120); only the figure of Neitsch *et al.* (2002) formulae was shown which illustrates the best performance among all other formulas. The results of correlation coefficient were shown in Figures 2 and 3.

As it was shown in figures 2 and 3 and table 3, according to Neitsch *et al.* (2002) formulae the lowest and the highest correlation coefficient in CNI, i.e., 997 and 0.999 were belonged to August and July and those for CNIII, i.e. 0.972 and 0.998 were matched to August and February, respectively. In addition,

CNI had much higher correlation coefficient with observed data but much lower average relative error and root mean square especially in July. The results of RE and RMSE were the same which established that Neitsch et al. (2002) formulae had the best performance in both CNI and CNIII and in CNI Chow et al. (1988) had the second rank. Sobhani (1975) and Hawkins et al. (1985) were ranked as the third and the fourth in terms of performance. However, in CNIII calculations, Sobhani (1975) showed the second rank and Hawkins et al. (1985) and Chow et al. (1988) were the third and the fourth, respectively. Finally, in both of them (CNI and CNIII) Mishra et al. (2008) was the weakest model in this study. These results could be resulted from the form of formula. The form of formulae in Neitsch et al. (2002) is exponential while in Mishra et al. (2008) and the rest of formulas are linear. Furthermore, this finding showed that the differences between values in this watershed were very high. These findings were in conformity with Deshmukh et al. (2013). In fact this area was situated in semiarid region and in this region there are much variability between rainfall and runoff in all over the year. Because precipitation in this area has not even distributed normally so that it may all rainfall amount fall in one place and in a short time with high intensity, therefore we may have a lot of errors. These results were agreed with Carless et al. (2011) and Lin et al. (2014) who revealed that this event may occur because of the greater kinetic energy of the rainfall which may cause more quickly alterations in the soil surface, surface seal forming and reduction in time of runoff. At last, increasing rainfall intensity and duration increase the total kinetic energy.

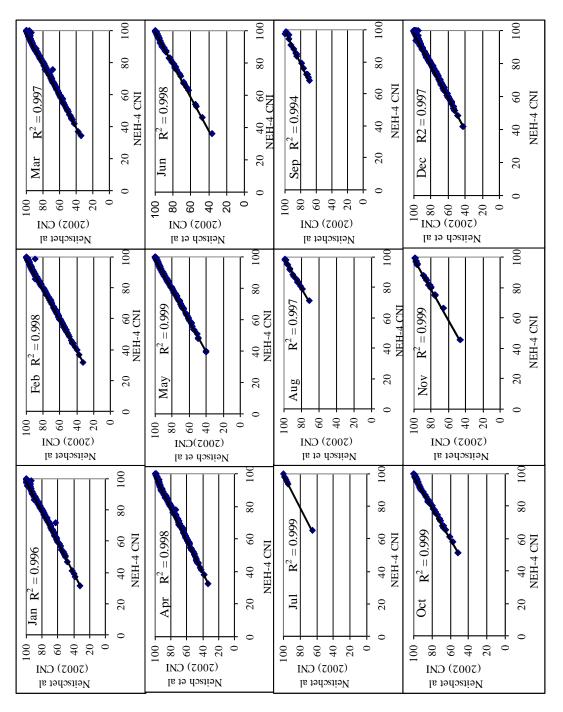


Figure 2 The performance of the best model (Neitsch et al., in CNI condition 2002)

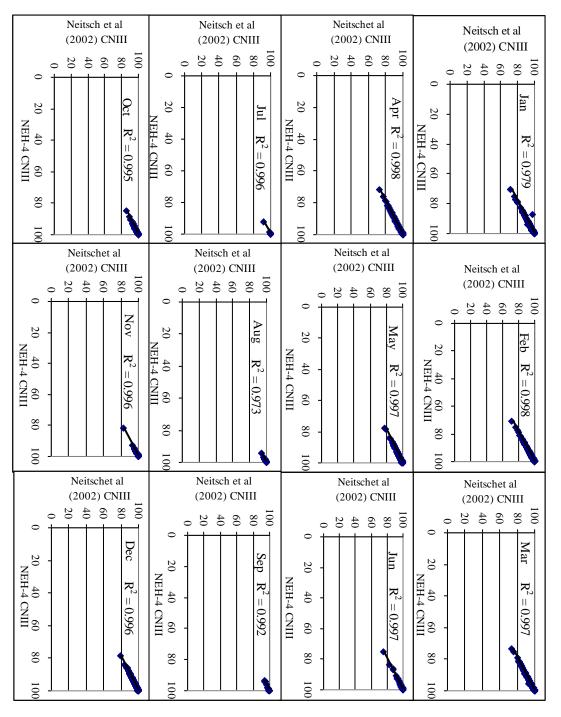


Figure 3 The performance of the best model (Neitsch et al., in CNIII condition 2002)

#### 5 CONCLUSION

This work tried to evaluate some methods which are related to runoff computation. One of the best models is SCS-CN which presented by Resources Conservation Services Natural (NRCS). It is a standard approach for runoff estimation in a watershed due to the simplicity and high speed of computing and applicability for agriculture and urban watershed. Some scientists have presented some methods such as Sobhani, Hawkins, Chow, Neitsch and Mishra. Therefore, in this article these methods were compared with each other and the results showed that Neitsch model was the best and Mishra was the poorest model in CNI and CNIII. Since the SCS-CN is used for estimation of storm runoff and it cannot be used for snow and base flow estimation, application of modified CN methods in other Watersheds are also advised. At the end, we recommended these matters should be investigated by using some techniques such as GIS and RS and in other watersheds with longer duration. In addition, land-use change should be considered, especially for its impact in the flood season. The role of land-use change should be appropriately considered due to its impact on water resources and ecosystem health in the watershed.

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## كارائي مدلهاي مختلف تخمين شماره منحني (مطالعه موردی: حوزه آبخیز بار در استان خراسان رضوی، ایران)

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چکیده از میان مدلهای موجود برای تخمین رواناب در مدیریت حوزه آبخیز، روش شماره منحنی سرویس حفاطت خاک همراه با تغییرات آن بهطور گستردهای برای سیستمهای آبخیز غیر تجهیز شده بهدلیل برآورد سریعتر و دقیقی از رواناب سطحی، در بین مدلهای دیگر می باشد. این دیدگاه به طور وسیعی توسط هیدرولوژیستها، متخصصان منابع آب، جنگلداران و مهندسین مورد قبول قرار گرفته است. بنابراین، مطالعه حاضر بهمنظور برآورد شماره منحنی با استفاده از مقادیر شماره منحنی از طریق روشهای مختلف از جمله Hawkins (۱۹۷۵) Sobhani ،SCS و همکاران (۱۹۸۵)، Chow و همکاران (۱۹۸۸)، Neitsch و همکاران (۲۰۰۸) و Mishra و همکاران (۲۰۰۸) در حوزه آبخیز بار در ایران انجام شد. نتایج نشان داد که فرمول Neitsch دارای بهترین نتایج در تبدیل شماره منحنی برای شرایط با رطوبت پیشین کم (CNI) و شماره منحنی برای شرایط با رطوبت پیشین زیاد (CNIII) بر اساس حداقل خطای نسبی بوده است. در حالی که ضعیفترین عملکرد مربوط به فرمول و روش Mishra و همکاران (۲۰۰۸) بوده است. وجود چنین نتایجی می تواند به علت شکل خاص در روش Neitsch و همکاران (۲۰۰۲) به دلیل حالت نمایی و هم چنین شرایط آب و هوایی حوزه آبخیز بار در طول یک سال باشد.

كلمات كليدي: تخمين سيلاب، رطوبت پيشين خاك، شمال شرقي ايران، مدلسازي بارش-رواناب