

Investigating the Relationship between Meteorological and Agricultural Droughts in Northwest Iran

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Aims: This study aims to assess the variations in drought trends using indices in Northwest Iran, as vegetative cover plays a vital role in environmental stability.

Materials & Methods: To achieve this goal, the study includes three stages: determining the Standardized Precipitation Evapotranspiration Index (SPEI) using monthly temperature and precipitation data from meteorological stations, calculating the Vegetation Health Index (VHI) based on derived datasets from MODIS satellite images for the period 2001-2021, and examining the correlation between indices to determine the duration of vegetation cover response to water scarcity and identify trends at 3, 6, 9, and 12-month time scales.

Findings: Based on the results of the Mann-Kendall test, the stable (48.56%) and increasing (50.43%) trends cover most of the studied areas, and a smaller area had a decreasing trend (1.01%). Additionally, positive correlations between VHI and SPEI were observed across all time scales. The SPEI-3 months showed the highest Pearson correlation (R²= 0.83) with VHI values for the growing season, indicating that water accumulation in the past three months significantly impacted vegetation cover.

Conclusion: This study emphasizes the necessity of monitoring and managing drought, focusing on vegetation cover status in the Northwest of Iran, especially in East Azerbaijan Provinces. It also introduces drought indices as a crucial component of the drought monitoring system.

Keywords: Trend Analysis, Drought Index, Time Scale, Northwest of Iran.

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Introduction

Vegetative cover is vital in soil conservation, climate regulation, hydrological processes, carbon cycles, and ecosystem stability [1]. It reflects the dynamic characteristics of terrestrial ecosystems. It is considered a sensitive indicator of ecosystems' response to climate change and human activities,
fundamentally indicating the overall fundamentally indicating the overall state of the ecological environment $[2, 3, 4]$. Spatial-temporal changes in vegetation cover patterns alter the regional landscape and impact ecosystem structure, leading to weaker resilience in maintaining and enhancing ecosystem sustainability [5]. Therefore, understanding the impact of water availability on vegetation cover growth in connection with climate changes is essential for predicting the dynamicity of vegetation cover against environmental changes. It also aids in developing effective policies and strategies for environmental protection and sustainability.

One of the climatic events affecting vegetation cover is vegetation drought. Recurring droughts are a primary cause of rangeland degradation and a significant factor in reducing agricultural crop yields, transforming forested lands into shrubdominated landscapes [6]. However, since vegetation cover exhibits specific resistance to drought, the effect of drought on vegetation growth has a distinct time lag $[7]$. Vegetation growth affected by previous drought is called the lag time effect. For instance, a drought that occurred one month, three months, or even twelve months ago stresses vegetation growth. In other words, vegetation growth may not be primarily due to current weather conditions, but the initial climatic conditions significantly impact vegetation growth. Therefore, considering the effects of time lag is crucial when investigating the mechanisms of the interactive effects of climate and vegetation cover $^{[8]}$.

Several indices, such as SPI $^{[9]}$ and SPEI
 $^{[10]}$ have been introduced to assess have been introduced to assess drought based on meteorological station precipitation data. However, these methods are effective only for evaluating drought conditions near meteorological stations due to limited spatial coverage, low station density, and uneven distribution of stations, especially in semi-arid regions, where precipitation varies significantly over short distances $[11]$. In contrast, remote sensing (RS) data become the primary data source for defining drought indices due to their extensive spatial coverage and high public accessibility [12]. Drought assessment indices based on remote sensing data include VHI $^{[13]}$, NDVI $^{[14]}$, VCI, TCI $^{[15]}$, and EVI $^{[16, 17]}$. Researchers have utilized various drought indices in their studies over the years. Potop et al. [18] analyzed spatial characteristics of drought and their trends at different time scales based on the SPEI index in the Czech Republic using data from 184 stations. These researchers reported a negative trend for most stations, indicating a stable or increasing trend for drought events. Tirivarombo *et al.* [19] compared SPI and SPEI in the Kafue River basin in northern Zambia. The results showed that temperature change is crucial in determining drought phenomena. Bazrafshan et al. [20] compared SPI, SPEI, and RDI indices in analyzing drought trends in dry and semi-dry regions of Iran for warm and dry climates, reporting an increasing trend (below -2.61) for most regions and a decreasing trend (above 2.61+) for cold and dry climates. Hosseini Pezhuh *et al.* ^[21] used the standardized precipitation index (SPI) and Standardized
Evapotranspiration Precipitation Index Evapotranspiration Precipitation (SPEI) for drought detection and evaluation in the Kesilan watershed. They reported that SPEI performed better than SPI, especially at the end of the spring and summer seasons. Wang *et al.* ^[22] focused on China in a study

aiming to investigate spatial and temporal developments in vegetation drought. They used the Vegetation Health Index (VHI) to identify vegetation drought and the Pixelbased Trend Identification Method (PTIM) to examine its spatial and temporal evolution from 1982 to 2020. The results indicated a decreasing trend in vegetation drought during the study period, with the most prominent reduction observed in spring. Other researchers have also reported a high correlation between VHI and vegetation cover performance, especially during critical growth stages $^{[23, 24]}$.

Different researches show that changes in climate affect plant growth. Therefore, it is essential to understand how climate influences vegetation cover across different environmental conditions. Since the impact of climate on vegetation cover has not been investigated in the northwestern Provincess of Iran, this study specifically focuses on meteorological and agricultural droughts, aiming to comprehend the occurrence trend of drought in Northwest Iran, as these two types of droughts are more directly related to the examination of climatic factors and vegetation cover performance.

Material & Methods Study Area

The northwest of Iran has different climate regimes and air masses, so diverse climate zones appear in its territory $[25]$. In this study, the Northwest of Iran encompasses the Provincess of East Azerbaijan, West Azerbaijan, Zanjan, and Kurdistan. The elevation gradually increases from east to west and decreases around Lake Urmia. The climate transitions from semi-arid to arid from north to south, and the temperature gradually rises from north to south. Due to its topographical features, location, and elevation, this region exhibits environmental, plant, and animal diversity adapted to the current climate conditions. The region's longterm fluctuation and variability in climatic factors can have significant environmental, agricultural, economic, and even human consequences [25]. The primary land covers in the study area include grasslands (56.81%), deserts (32.90%), agricultural lands (4.78%), water bodies (2.18%), and sparse shrublands (1.68%) (Figure 1).

Data collection

To calculate the SPEI, monthly temperature and precipitation data from meteorological stations in Ardabil, Khorram Dareh, Khoy, Maragheh, Sanandaj, Tabriz, Urmia, and Zanjan were utilized. These selected stations were chosen due to their common statistical period, location within the study area, or proximity to the study region. The geographical coordinates of each station are provided in Table 1. For the Vegetation Health Index (VHI), data from the MOD13Q1 sensor of the Moderate Resolution Imaging Spectroradiometer (MODIS) for 2000 to 2023 were employed.

Estimation of Drought Indices

Drought indices are essential for evaluating drought severity, modeling, and predicting drought [27]. Each of these indices is associated with one or more variables that impact environmental characteristics such as precipitation, temperature, river flow, soil moisture, and other variables [28]. Most drought indices are tailored to specific geographical areas and objectives, making their selection for achieving accurate and comprehensive analyses challenging due to the inherent complexities of the phenomenon $[29]$. To overcome this challenge, considering the enhanced role of temperature in drought propagation due to climate change $[10]$, the SPEI was employed. Additionally, the Vegetation Health Index (VHI) was chosen among satellite-based indices due to its consideration of local biophysical factors (soil and slope) and climatic conditions $[30, 31]$.

Figure 1) Geographical location of the study area in Iran.

Estimation of Standardized Precipitation Evapotranspiration Index (SPEI)

The SPEI is employed to quantify the intensity and duration of drought events [32]. SPEI is based on the difference between precipitation (water supply) and potential evapotranspiration (water demand), making it applicable for assessing meteorological and hydrological droughts [10]. Drought indices based on meteorological data benefit from the advantage of long-term temporal scales [33]. SPEI is scalable for time intervals ranging from 1 to 48 months or more. SPEI-1 represents drought conditions in the current month, while SPEI-12 reflects the overall annual drought conditions. The multiscale nature of SPEI enables the identification of various types of droughts and their impacts on different systems $[32]$.

The SPEI equation was processed for four time scales (3, 6, 9, and 12 months) using functions defined in MATLAB. MATLAB is a computational software platform that encompasses over 600 mathematical operation functions extensively utilized in climate analysis projects $[34]$.

Estimation of vegetation health index (VHI) VHI has been proposed to quantify vegetation drought using several years of vegetation water-deficit situations [35]. VHI is a linear combination of the Vegetation Condition Index (VCI) and the Temperature
Condition Index (TCI), simultaneously Condition Index (TCI), reflecting vegetation cover's greenness and temperature status. The VHI was computed according to the Eqs. (1) , (2) , and (3) using ArcMap 10.8.2 ^[30].

$$
VHI = 0.5 \times (VCI) + 0.5 \times (TCI) \qquad Eq. (1)
$$

$$
VCI = \frac{EVI_i - EVI_{\min}}{EVI_{\max} - EVI_{\min}} \times 100
$$
 Eq. (2)

$$
TCI = \frac{LST_i - LST_{\min}}{LST_{\max} + LST_{\min}} \times 100
$$
 Eq. (3)

In the provided equations, *EVI i* represents the EVI value of a pixel in year *i*, while *EVImax* and EVI_{min} denote the maximum and minimum values of EVI for the same pixel from 2001 to 2019, respectively. Moreover, LST stands for Land Surface Temperature, where LST_i

signifies the LST value of a pixel in the year I, and *LST*_{max} and *LST*_{min} indicate the maximum and minimum values of LST for the same pixel from 2001 to 2019, respectively.

Trend analysis

Trend determination tests are divided into two types: parametric and nonparametric. The trend test applied in this study is the nonparametric Mann-Kendall (MK) test. The Mann-Kendall test is effectively used to detect natural fluctuations or trends in the changes of a process. This rank-based test is suitable for detecting non-linear trends $[36]$. The advantage of the nonparametric test, such as the Mann-Kendall test, over parametric tests like the t-test is its suitability for time series that do not follow a normal distribution and for datasets with missing or deleted values [37]. The Mann-Kendall test is calculated based on the Eq. (4).

$$
S = \sum_{i=2}^{n} \sum_{j=1}^{i-1} (sign(xi - xj))
$$
 Eq. (4)

In which xi and *xj* represent consecutive data, n denotes the length of the time series, and the sign function is computable as Eq. (5):

$$
sign(xi - xj) = \begin{cases} +1 & \text{if } (xi - xj) > 0 \to \\ 0 & \text{if } (xi - xj) = 0 \\ -1 & \text{if } (xi - xj) < 0 \end{cases}
$$

 $Eq. (5)$

The mean E(S) and variance Var(S) of the statistic *S* were calculated as Eq. (6) and Eq. (7):

$$
E(S) = 0
$$
 Eq. (6)

$$
VAR(S) = \frac{1}{18} [n(n-1)(2n+5)
$$

- $\sum_{p=1}^{q} tp(tp-1)(2tp+5)]$ Eq. (7)

Where *tp* is the number of sequences for

the p^{th} value, and p is the number of values in the sequences. The second part in the above equation represents an adjustment for sensitive sequences or data. The standardized statistic *ZM* for the test is obtained from the Eq. (8):

$$
Z_M = \begin{cases} \frac{S-1}{\sqrt{VAR(S)}} & \text{if } S > 0 \rightarrow \text{increasing trend} \\ 0 & \text{if } S = 0 \rightarrow \text{without trend} \\ \frac{S+1}{\sqrt{VAR(S)}} & \text{if } S > 0 \rightarrow \text{decreasing trend} \end{cases}
$$

Eq. (8)

The positive value of the *ZM* trend indicates an increasing trend, while the negative value of *ZM* indicates a decreasing trend in the time series ^[38]. The threshold of statistical significance at a five percent level is |*ZM*| > 1.96. The Earth Trends Modeler (ETM) in Terrset software was employed to conduct this test.

Additionally, the correlation of VHI concerning multitemporal SPEI was determined using the coefficient of determination based on Eq. (9).

$$
R^2 = 1 - \frac{RSS}{TSS}
$$
 Eq. (9)

In the given equation, R^2 stands for the coefficient of determination. *RSS* and *TSS* represent the Sum of Squares of Residuals and the Total Sum of Squares, respectively. Higher values of these parameters mean more substantial agreement between the compared data, as elucidated by Pogačar *et al.* ^[39].

Moreover, the slope coefficient of the VHI changes concerning multitemporal SPEI was ascertained using Eq. (10).

$$
Slope = \frac{n \sum_{n=1}^{i=1} x_i y_i - \sum_{n=1}^{i=1} x_i \sum_{n=1}^{i=1} y_i}{n \sum_{n=1}^{i=1} x_i^2 - (\sum_{n=1}^{i=1} x_i)^2}
$$
 Eq. (10)

In the provided equation, X_i and Y_i denote the values of the independent and dependent

variables in the *i*th year, respectively. The *n* represents the years during the study period (*n=20*).

The negative slope value indicates that the dependent variable exhibits an increasing trend, while a positive slope value signifies a decreasing trend in the dependent variable.

Findings

Maximum vegetation cover

According to the map presented in Figure 2, the highest percentage of vegetation cover during the study period is observed in April, while the lowest is in June.

Figure 2) Map of vegetation picked in different months.

Vegetation trend

Figure 3 illustrates the trend of vegetation changes in the study areas. According to the map, yellow pixels indicate a "no-change" trend, green pixels represent an "increasing trend" (drying trend), and red pixels indicate a "decreasing trend" or relief from drought conditions. The stable (48.56%) and increasing (50.43%) trends cover most of the studied areas, while a few pixels have been allocated decreasing trend (1.01%) conditions.

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Figure 3) Map of vegetation trend over time.

The results of the Mann-Kendall test are presented in Table 2. According to the provided information, East Azerbaijan Provinces exhibits the highest increasing trend (5.68), while Kurdistan Provinces shows the lowest (4.44). Moreover, on average, the maximum and minimum changes in the vegetation cover trend are observed in East Azerbaijan (1.79) and Kurdistan (1.02). Also, the Standard Deviation (SD) shows that the most changes occur in West Azerbaijan (1.01), while the most minor change occurs in Zanjan (0.87), which represent the heterogeneous and homogeneous changes, respectively, in these Provincess.

Figure 4) Maps of correlation between VHI and Multitemporal SPEI.

Figure 5) Correlation between VHI and multitemporal SPEI in different Provincess

Correlation between VHI and Multitemporal SPEI

The temporal response of vegetation cover to drought and the spatial distribution of correlation between VHI and multitemporal SPEI is illustrated in Figure 4. A significant positive correlation exists across the entire study area for all temporal scales. The spatial
distribution of correlation coefficients distribution of correlation gradually increases towards the southeast for all studied temporal scales. The lowest correlation values are observed in the 6 and 9-month scales. However, (SPEI-3),

Figure 6) Maps of the absolute value of slope in the linear regression VHI and multitemporal SPEI

calculated for the growing season, exhibits the highest Pearson correlation with VHI values (R^2 = 0.83). In this particular scale, the northeastern and southwestern parts of the study area are identified as the most sensitive regions to drought conditions, being more prone to drier conditions (e.g., drought).

As shown in Figure 5, the correlation value decreases when the time scale increases. The highest correlation of the studied indices is observed at the 3-month scale, attributed respectively to the Provincess of Kurdistan, Ardabil, West Azerbaijan, East Azerbaijan, and Zanjan. Consequently, the 3-month scale is more suitable for determining the intensity and duration of vegetation-related drought.

Table 2) Results of the MK test.

Linear regression between VHI and multitemporal SPEI

Figure 6 illustrates the maps of the absolute value of the linear regression slope between VHI and Multitemporal SPEI. It is observed that as the time scale increases, the absolute value of the slope in the linear regression between VHI and multitemporal SPEI decreases. The highest and lowest slopes of changes are attributed to the 3-month and 9-month scales, respectively.

According to the presented data in Figure 7, the highest magnitude of slope changes is observed in the 3-month scale, respectively, belonging to the Provincess of Kurdistan, East Azerbaijan, Ardabil, West Azerbaijan, and Zanjan.

Figure 7) The absolute value of slope in the linear regression between VHI and multitemporal SPEI in different Provincess

Discussion

The highest and lowest percentages of vegetation cover were observed in East Azerbaijan and Kurdistan Provincess, respectively, during the study period. These results emphasize the importance of vegetation mapping or classification in assessing natural environments, as it provides valuable information by quantifying vegetation cover over a specific period. Such information is crucial for natural resource management and the implementation of environmental conservation and restoration $programs$ ^[40,41].

Changes in drought trends, including no change, stability, and increasing trends, were evident across the study area. An increasing trend indicates a decline in vegetation cover over time, suggesting that drought is a recurring and growing phenomenon in the region. Some researchers have identified global warming $[42]$, regional land use $[43]$, and land cover changes $[44]$ as potential drivers of increasing trends. Moreover, some studies have highlighted the role of global warming in assessing the sustainability of wheat and barley farming systems in the Maragheh-Bonab plain of East Azerbaijan Provinces [45]. Although it is challenging to attribute the identified trends to specific factors accurately,

temperature variations have been introduced as a fundamental factor influencing climate variability in arid and semi-arid regions like Iran $[46]$. Previous researchers have
proposed potential evapotranspiration, evapotranspiration, actual evapotranspiration, and temperature as primary factors in agricultural drought [47]. The results of this study illustrated a general increasing trend in Northwest Iran, consistent with Dehghani Sargazi *et al.* [48]. The increasing trend in East Azerbaijan Provinces, which has the largest area in the Northwest of Iran, may exacerbate the conditions of agricultural lands in this Provinces. Some research sources have indicated that 46% of agricultural lands in East Azerbaijan Provinces are located in the Lake Urmia basin, and this lake is at risk of drying up due to the expansion of agricultural activities in its watershed and the lack of natural water rights [49]. Therefore, it is essential to develop a plan to reduce the potential effects of drought on agricultural production by considering a combination of climate variables, including precipitation and temperature.
The significant

significant positive correlation observed across the entire region for all studied time scales was confirmed. It is crucial to mention that the effectiveness of SPEI in assessing vegetation drought and, consequently, biomass production has been validated by previous studies $[10]$. Similarly, the effectiveness of VHI in understanding various aspects of vegetation cover [10, 24] and monitoring the onset of vegetation drought as an early warning system has also been confirmed $[50]$, reinforcing the validity of the results in this section.

Raja and Gopikrishnan^[51] compared and analyzed the performance of SPEI at time scales of 1, 3, 6, 9, and 12 months of temporal and spatial variation at 12 weather stations in Barmer arid region from 1979 to 2013. A modified MK test is used to determine the significance of drought-characteristic trends. The results showed that with the increase of the time scale, the temporal changes in SPEI became more consistent. The MK test showed that SPEI showed decreasing trends in 1and 3-month scales but increasing trends in 3, 6, and 12-month scales.

In this case, Musei *et al.* ^[52] used SPEI in 1, 3, 6, and 12-month time scales to evaluate the spatio-temporal changes of drought occurrence in Somalia. Temporal changes in drought showed a decreasing trend in severity and an increasing trend in drought duration with increasing time scales of SPEI. The Northeast and southwest parts of the study area were the most sensitive (vulnerable) regions to drought. Hence, these areas should be prioritized in drought planning. The 3-month time lag of vegetation response to climatic conditions was observed. One reason is the presence of annual plants and the growth form of grasses, which mainly cover the region and are dependent on short-term moisture. In other words, in the 3-month scale, the influence of climate factors on vegetation cover is more pronounced, and vegetation responds more quickly to climate in this scale. The time delay in the response of vegetation to climatic variables has been reported in various studies. For instance, a time lag of about three months was reported for the Eurasian continent $[53]$. In contrast positive relationships between vegetation cover and rainfall in one month were reported in the southwestern United States [54], Australia inland [55-56], Yun-Gui Plateau, and Inner Mongolia $[57]$.

The highest slope of changes was observed at the 3-month scale. The areas with higher slope changes are more sensitive or vulnerable to drought. Based on this, it can be inferred that SPEI-3 indicates the dominant drought pattern in the northwest part of Iran.

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The results of this study contribute to a better understanding of the evolution of agricultural drought in Northwest Iran in spatial and temporal dimensions during 2001- 2020. Additionally, future development plans can be directed toward supporting regions prone to agricultural drought using the obtained information.

Conclusion

This study investigated the physical characteristics of agricultural drought in the Northwest of Iran. The results of this study indicate that the time lag in vegetation response to drought is three months. The necessity of paying attention to current water resource management and adopting long-term regional policies in East Azerbaijan Provinces was highlighted due to an increasing trend in drought. Correlation results between VHI and SPEI confirmed the existence of a significant positive correlation throughout the region for all time scales, especially with 3-month time scales. These results highlighted the importance of implementing land use policies about climate change and considering the relationships between climate (processes that affect vegetation cover) and vegetation. However, considering other biophysical factors, such as soil characteristics and land characteristics, and paying attention to groundwater in the analytical framework can improve the accuracy and usefulness of the results. Also, considering that drought is the dominant phenomenon limiting vegetation growth in arid and semi-arid areas, calculating the slope of changes between indicators helps predict the occurrence of drought in the future.

The exposure of the northeastern and southwestern parts of the study area to drought and the need for measures to reduce the effects of drought were also reported. Overall, the findings of this research can serve as scientific evidence for predicting future drought trends, implementing
optimal management strategies, and optimal management strategies, and enhancing preparedness and resilience against drought.

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